

An inverse problem of magnetization in geoscience

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February 26, 2015



Context



L. Baratchart, J. Leblond and D. Ponomarev
(APICS team, Inria Sophia-Antipolis, France),



E. Lima and B. Weiss
(Earth, Atmospheric and Planetary Sciences Dpt., MIT,
Cambridge Massachusetts, USA)



and **D. Hardin and E. Saff**
(Center for Constructive Approximation,
Vanderbilt University, Nashville, Tennessee, USA).

- ▶ Geophysicists at MIT: study the story of Earth's magnetic field.
 ↪ by analysing magnetization characteristics of rocks.
- ▶ Not directly observable ↪ one observes the induced magnetic field.

Outline

Motivation

Strategies

Preliminary results

Why study planetary magnetic field?

- ▶ Magnetic field is useful:
 - ▶ For navigation (compass, migratory birds, some fishes, etc.).
 - ▶ It prevents stripping of the atmosphere by the solar wind.
- ▶ Complex phenomenon: generated by a “dynamo”.
 - ↪ Several possible mechanisms. Still fairly misunderstood.
- ▶ Polarity reversals.
 - ↪ One of the most convincing evidence of continental drift.

Hot questions:

- ▶ Did the moon have a dynamo?
- ▶ If so, what was generating it?
- ▶ When did it turn on/off?
- ▶ same questions for Mars (could explain why Mars lost its atmosphere).

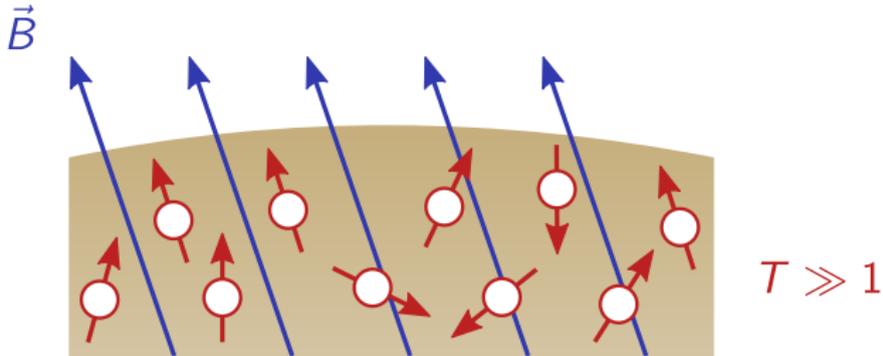
↪ A key question for understanding the early history of the solar system.

How do rocks acquire magnetization?

- ▶ Types of rocks: mainly
 - ▶ igneous (e.g., from volcanos);
 - ▶ or sedimentary (e.g., at the bottom of oceans).
- ▶ Thermoremanent magnetization:
ferro-magnetic particles follow the magnetic field.

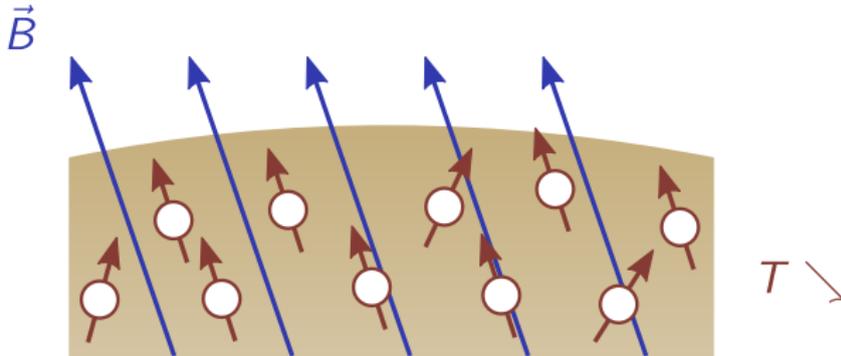
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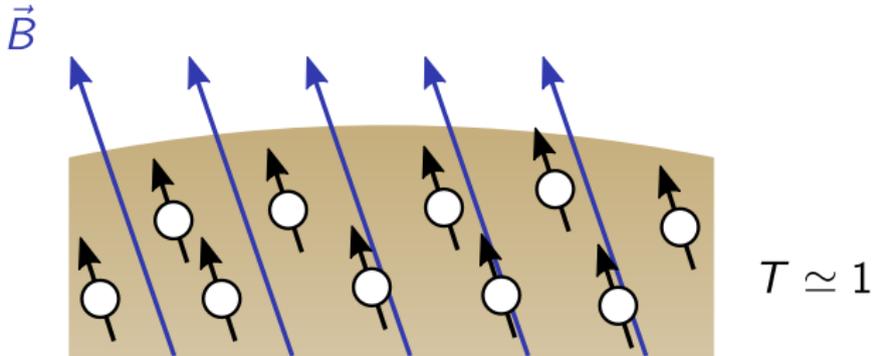
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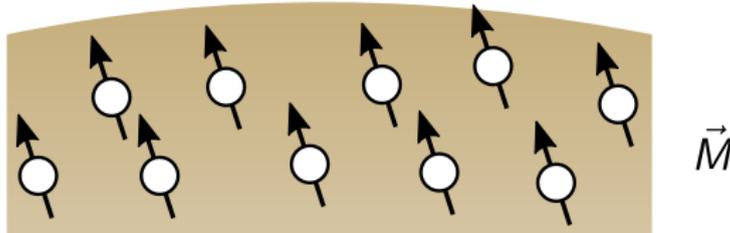
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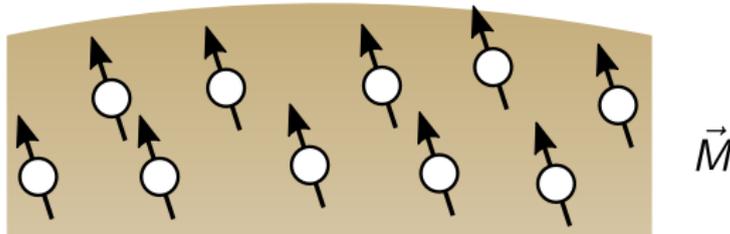
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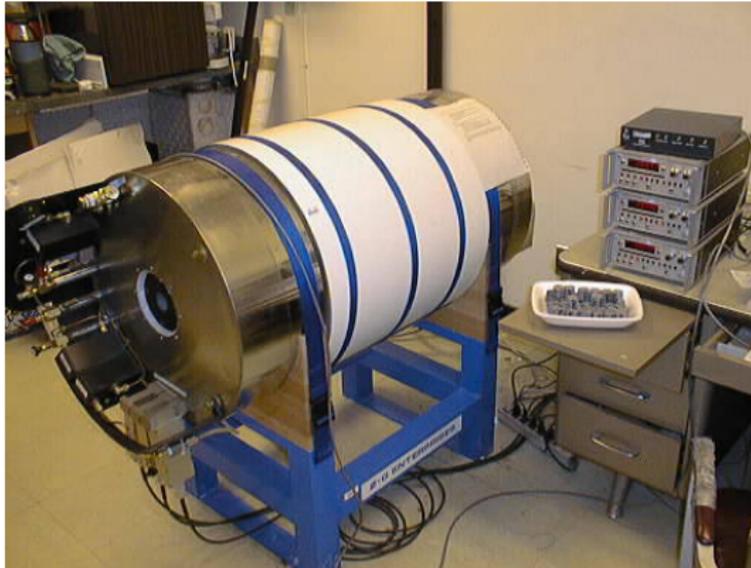
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- ▶ Can be subsequently altered
↪ under high pressure or temperature.

Measuring instruments

- ▶ Magnetometer: gives the net moment of a sample: $\iiint_{\text{rock}} \vec{M}$.



Measuring instruments

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- ▶ Scanning Magnetic Microscopes (SMM):

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- ▶ Scanning Magnetic Microscopes (SMM):

SQUID sensors

(Superconducting QUantum
Interference Device)

- ▶ high sensibility,
- ▶ far from the sample (100 μm),
- ▶ do not affect the magnetization,
- ▶ complicate to operate.

Non-superconducting sensor

- ▶ less sensitive,
- ▶ close to the sample (6 μm),
- ▶ may induce magnetizations,
- ▶ easy to operate.

SQUID microscope

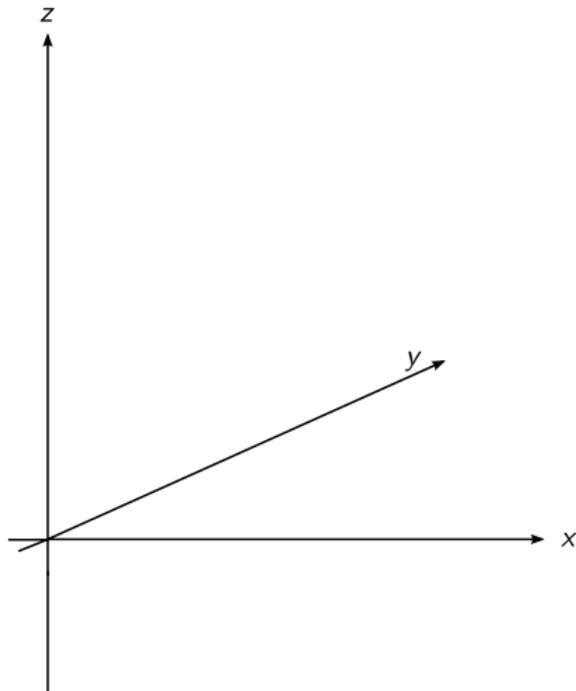


Pedestal + sensor

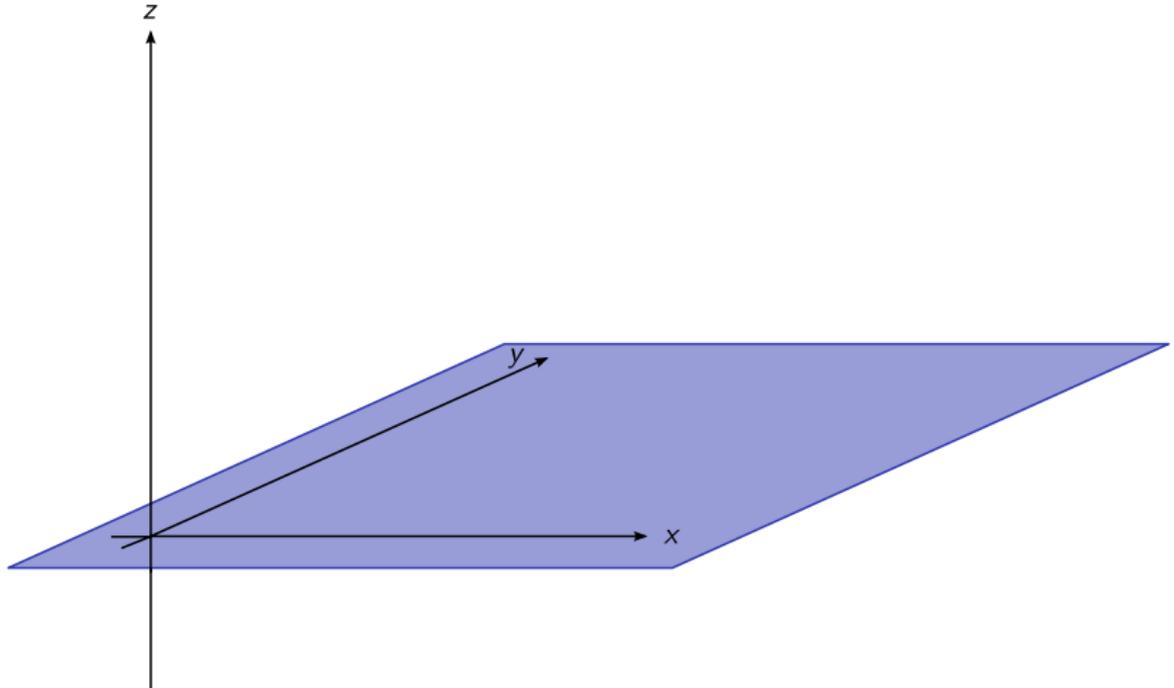


Sapphire window

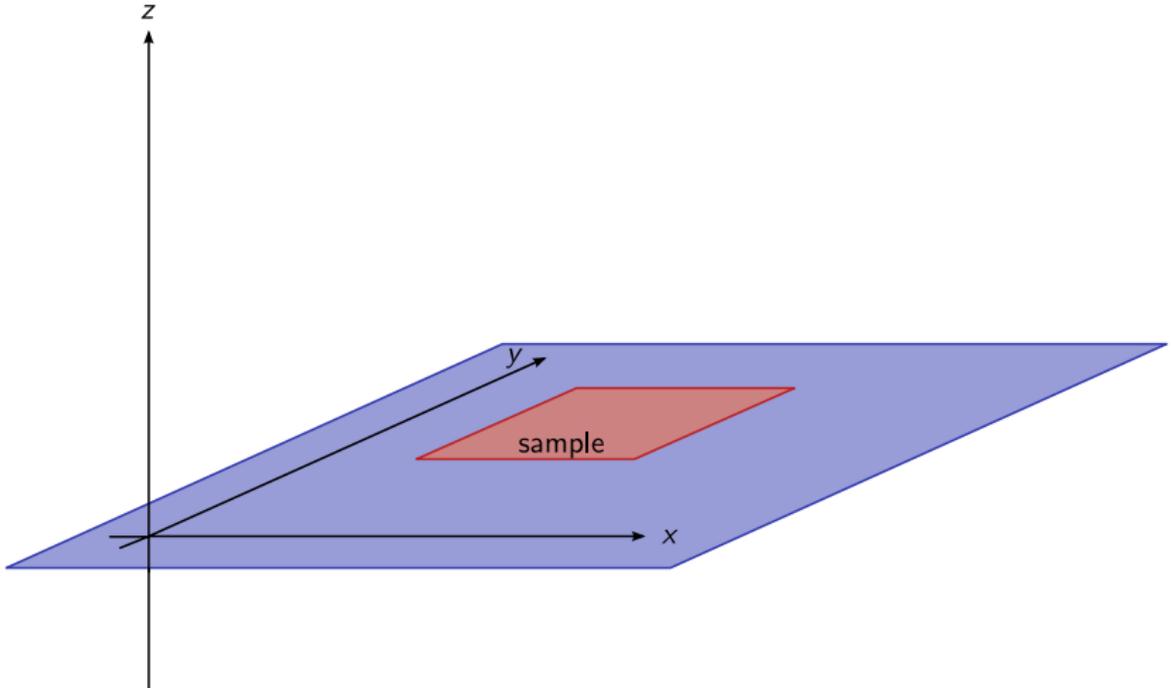
General scheme



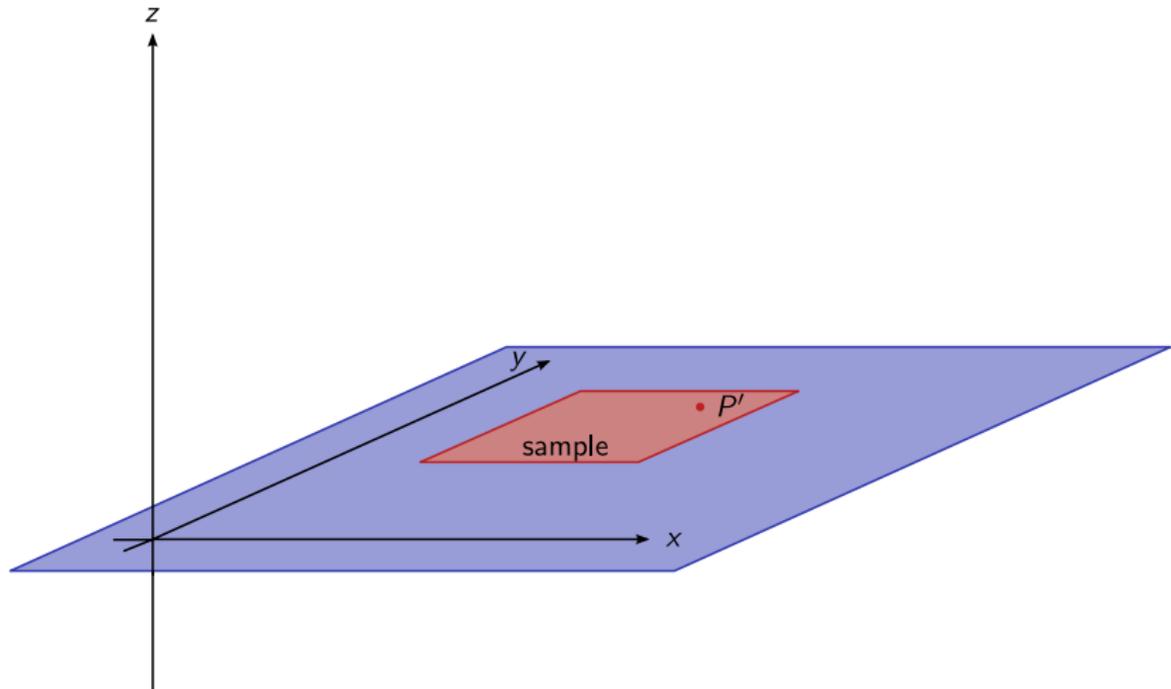
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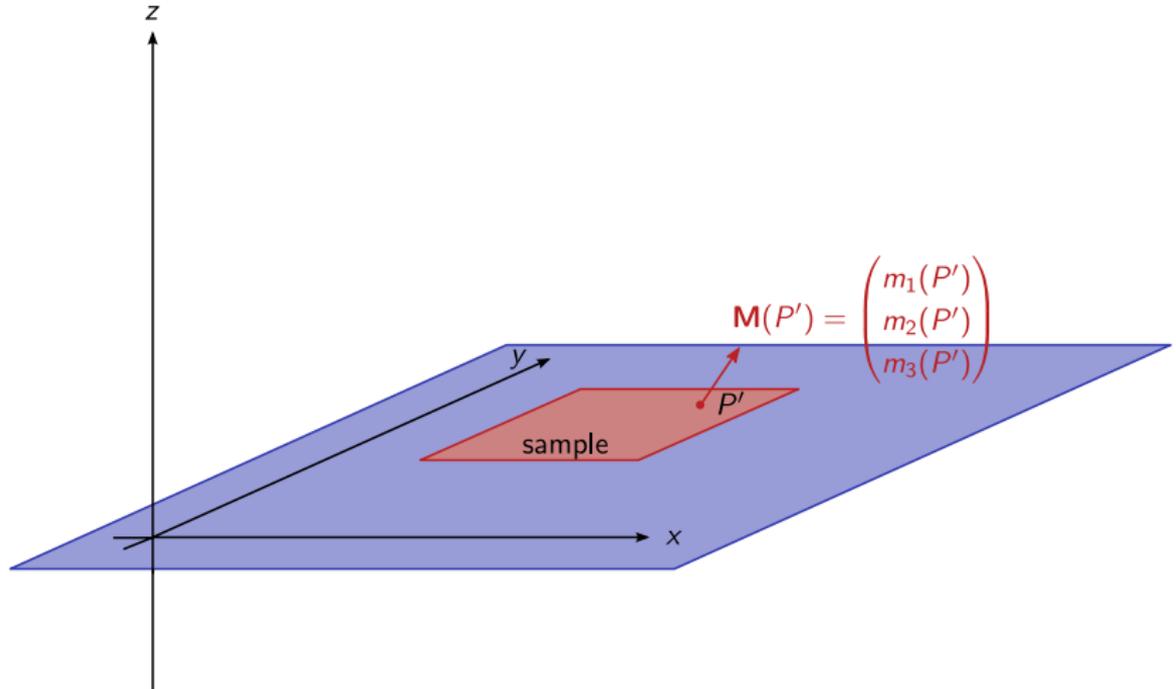
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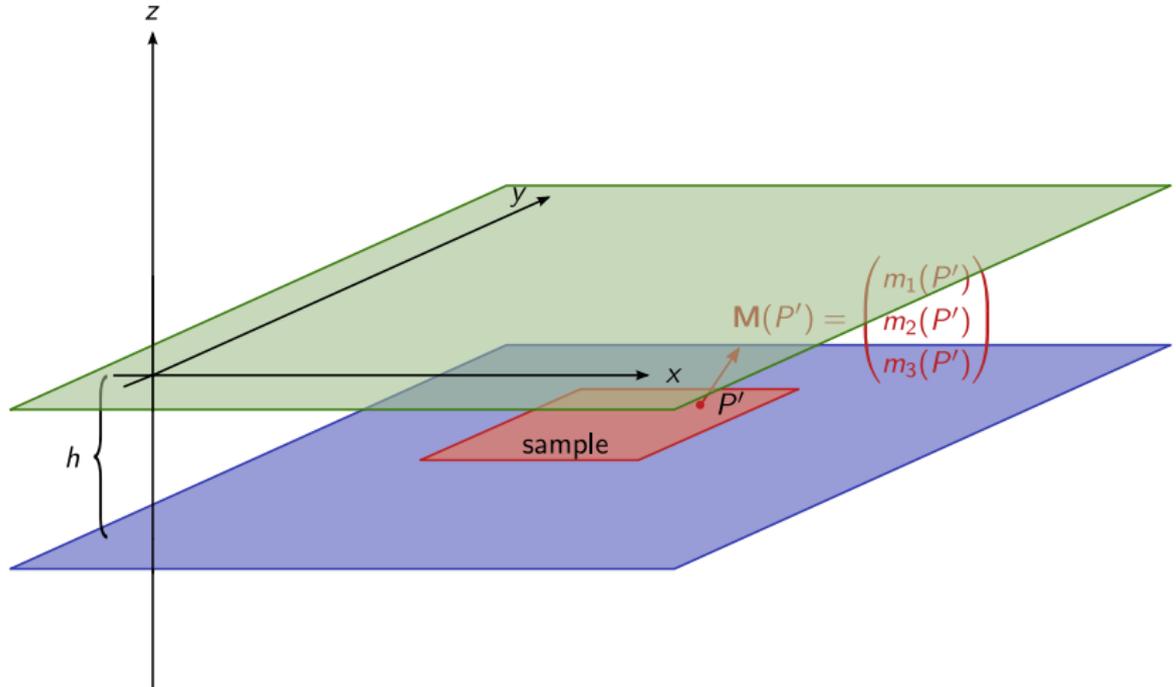
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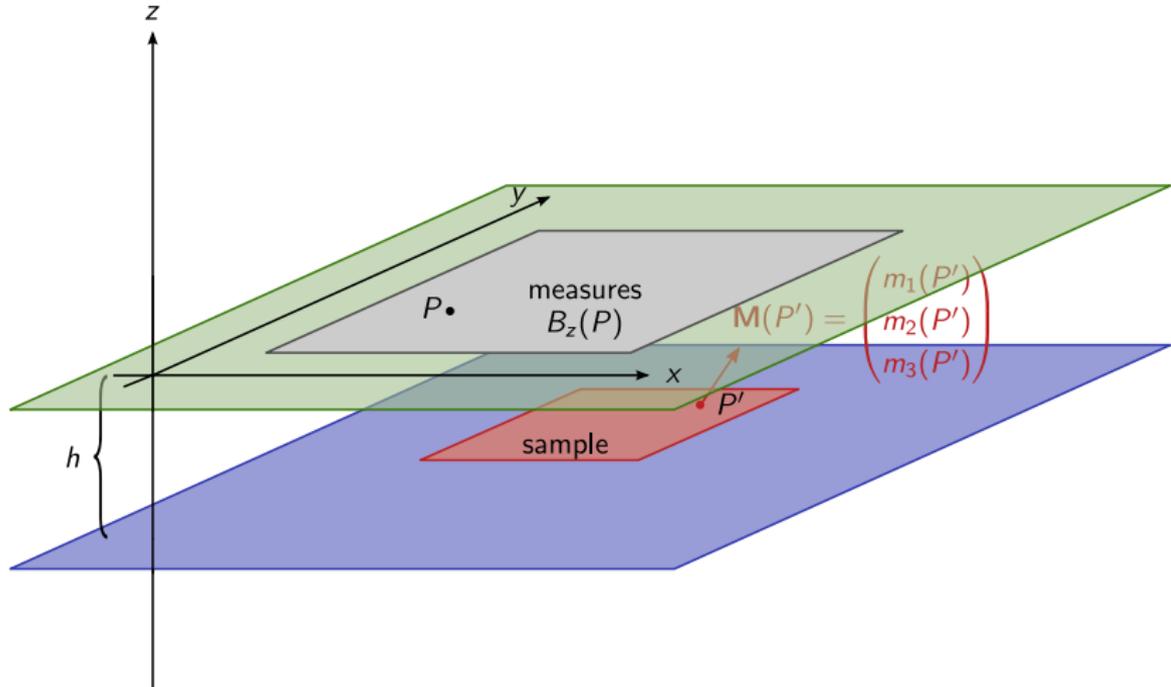
General scheme



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General scheme



Inverse problem

- ▶ Magnetization:

at each point $P' = \begin{pmatrix} x' \\ y' \\ z' \end{pmatrix}$ of the sample: $\mathbf{M}(P') = \begin{pmatrix} m_1(P') \\ m_2(P') \\ m_3(P') \end{pmatrix}$.

- ▶ Thin-plate hypothesis: $\mathbf{M} \neq 0$ only for $z' = 0$, i.e. $P' = \begin{pmatrix} x' \\ y' \\ 0 \end{pmatrix}$.

- ▶ Generates a magnetization potential:

at any point $P = \begin{pmatrix} x \\ y \\ z \end{pmatrix}$ of the space, $\varphi_M(P)$, s.t. $\Delta\varphi_M = \text{div}(\mathbf{M})$.

$$\varphi_M(P) = \frac{1}{4\pi} \iint \frac{m_1(P')(x - x') + m_2(P')(y - y') + m_3(P')z}{|P - P'|^3} dx' dy'$$

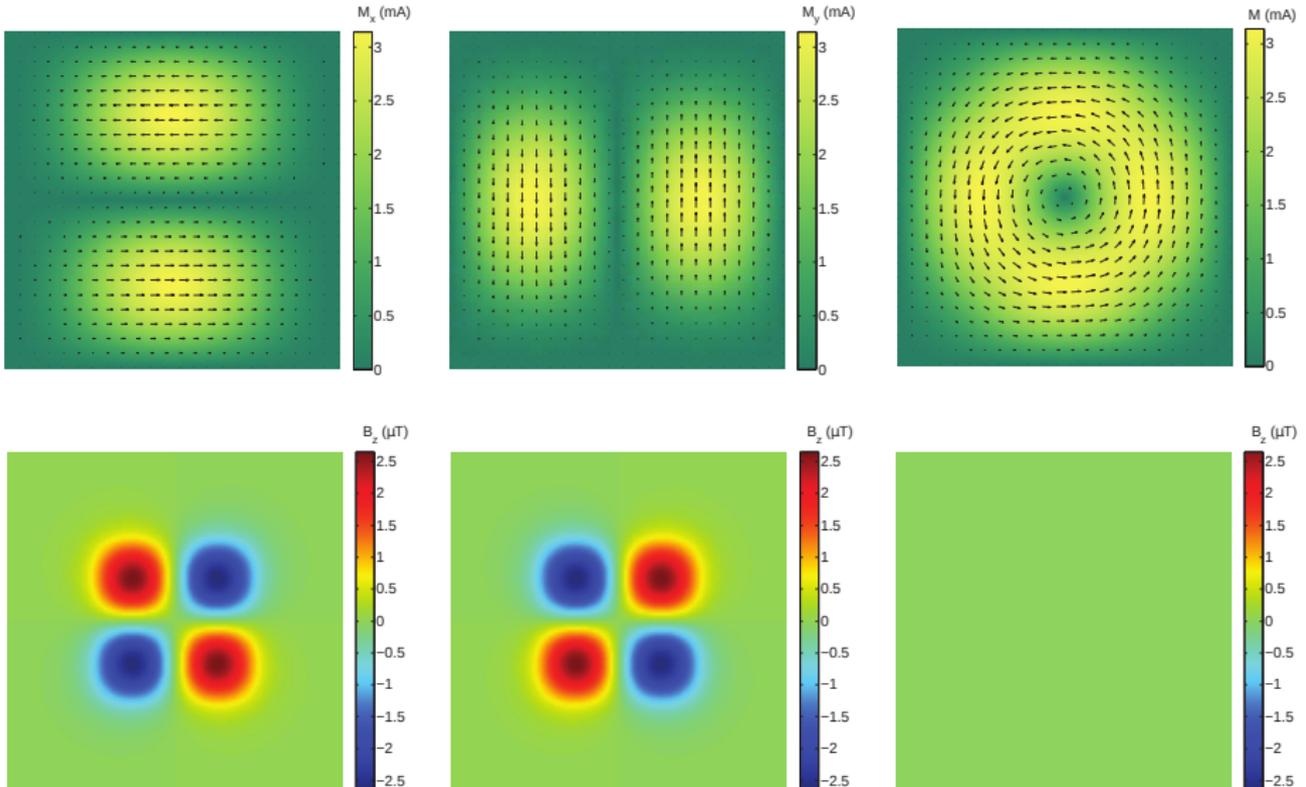
- ▶ Magnetic field induced by φ_M : $\mathbf{B}(P) = \nabla\varphi_M(x, y, z)$.

Silent sources

- ▶ **Inverse problem:** from measurements of \mathbf{B} at height h , recover \mathbf{M} (more precisely: only B_z is measured).
- ▶ \mathbf{M} is said **silent from above** (resp. below) if

$$\varphi_M(x, y, z) = 0 \text{ for all } z > 0 \text{ (resp. } z < 0).$$
- ▶ Such magnetizations exist. \rightsquigarrow Problem is ill-posed.

A silent source



Regularization hypotheses (1/2)

- ▶ Need for further assumptions on \mathbf{M} :
 - ▶ The support of \mathbf{M} is compact.

- ▶ Does compact support help? \rightsquigarrow **A bit**. In this case:
 - (\mathbf{M} is silent from above) \Leftrightarrow (\mathbf{M} is silent from below).

- ▶ There are silent magnetization from both sides: those s.t.

$$m_3 = 0 \text{ and } \operatorname{div} \begin{pmatrix} m_1 \\ m_2 \end{pmatrix} = 0.$$

Regularization hypotheses (2/2)

- ▶ \mathbf{M} can often be supposed **unidirectional**:

$$\exists \mathbf{v} \in \mathbb{R}^3, \exists Q : \mathbb{R}^2 \rightarrow \mathbb{R} \quad \text{s.t.} \quad \forall P', \mathbf{M}(P') = Q(P')\mathbf{v}.$$

\rightsquigarrow realistic if the rock has not been altered after its formation.

- ▶ Does unidirectionality help? \rightsquigarrow **No**.

For any \mathbf{M} and any direction \mathbf{v} not horizontal, there exists a scalar field Q s.t. $Q(P')\mathbf{v}$ is equivalent from above to \mathbf{M} .

- ▶ If we have compact support **and** unidirectionality? \rightsquigarrow **Yes**.

If \mathbf{M} is unidirectional and compactly supported, there is no unidirectional compactly supported equivalent magnetization.

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Fourier technique

- ▶ We consider Fourier transform with respect to horizontal variable $\mathbf{w} = \begin{pmatrix} x \\ y \end{pmatrix}$:

$$\hat{f}(\boldsymbol{\kappa}, z) = \iint f(\mathbf{w}, z) e^{-2i\pi(\mathbf{w} \cdot \boldsymbol{\kappa})} d\mathbf{w}.$$

- ▶ For the potential, we get at height $z > 0$:

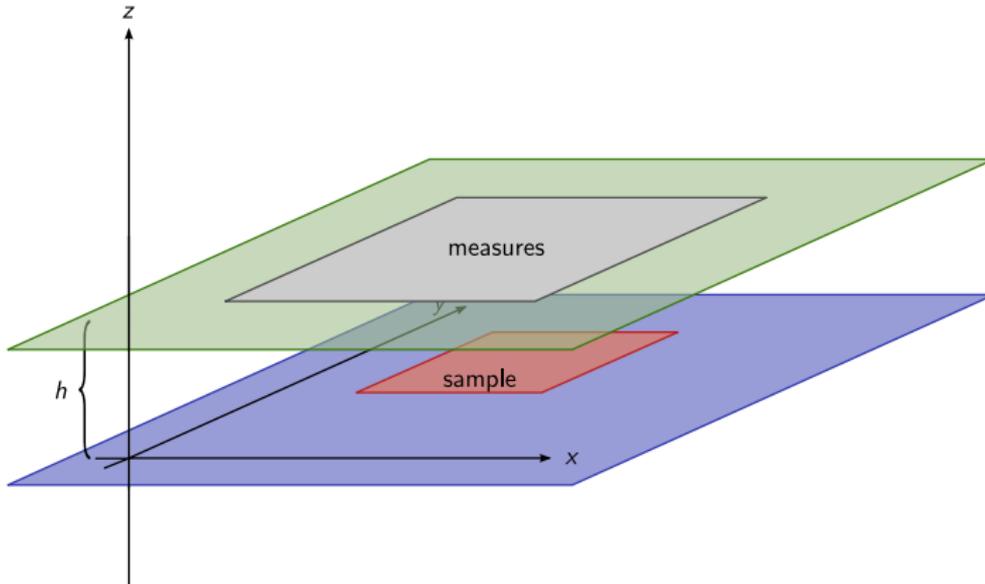
$$\hat{\varphi}_M(\boldsymbol{\kappa}, z) = \frac{e^{-2\pi z|\boldsymbol{\kappa}|}}{2} \left(i \frac{\boldsymbol{\kappa}}{|\boldsymbol{\kappa}|} \cdot \begin{pmatrix} \hat{m}_1(\boldsymbol{\kappa}) \\ \hat{m}_2(\boldsymbol{\kappa}) \end{pmatrix} - \hat{m}_3(\boldsymbol{\kappa}) \right).$$

- ▶ Case when \mathbf{M} is unidirectional: $\mathbf{M}(P') = Q(P')\mathbf{v}$:

$$\hat{\varphi}_M(\boldsymbol{\kappa}, z) = \frac{e^{-2\pi z|\boldsymbol{\kappa}|}}{2} \left(i \frac{\boldsymbol{\kappa}}{|\boldsymbol{\kappa}|} \cdot \begin{pmatrix} v_1 \\ v_2 \end{pmatrix} - v_3 \right) \hat{Q}(\boldsymbol{\kappa}).$$

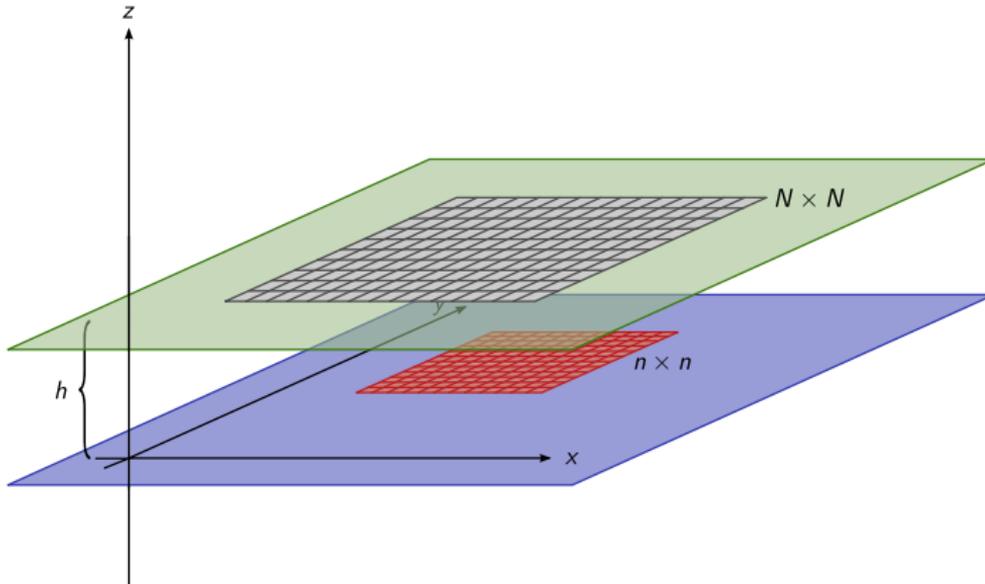
Direct approach

- ▶ Discretize the operator $\mathbf{M} \mapsto B_z$:
 - ▶ Rectangular regular $n \times n$ magnetization grid.
 - ▶ At each node of the grid a dipole.
 - ▶ Rectangular regular $N \times N$ measurement grid.



Direct approach

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Direct approach

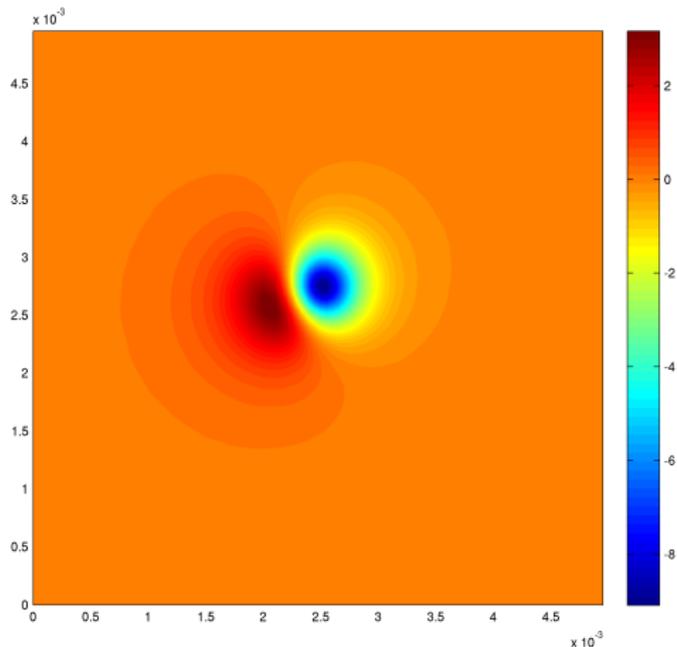
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 - ▶ At each node of the grid a dipole.
 - ▶ Rectangular regular $N \times N$ measurement grid.

- ▶ Let A be the matrix of the discrete operator.
 - \rightsquigarrow each column: field generated by a single dipole.

- ▶ Let b be the vector of measurements. *We try to solve $Ax \simeq b$.*

- ▶ Problem: A very big $\rightsquigarrow N^2$ rows and $3n^2$ columns
 - \rightsquigarrow example: $N = 100, n = 50 \rightsquigarrow 600$ MB just to store A .

Lunar spherule example



Field measured for the lunar spherule.

Direct approach (2)

If columns of A are linearly independent:

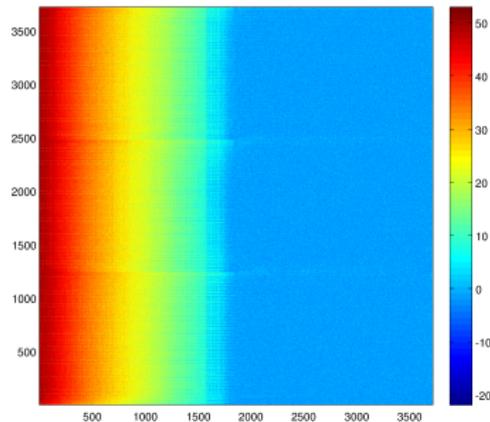
$$(\text{minimize } \|Ax - b\|_2) \Leftrightarrow (\text{solve } A^*Ax = A^*b).$$

- ▶ Interesting if $3n^2 \ll N^2$, i.e. magnetization is localized.
- ▶ Computing A^*A and A^*b : simple dot-products.
- ▶ SVD decomposition of a matrix $M = A^*A$: $M = VSV^*$ where:
 - ▶ S is diagonal and non-negative.
 - ▶ V is orthogonal.
- ▶ $x = VT$ where $T = S^{-1}V^*A^*b$. In other words $x = \sum_k t_k V_k$.
- ▶ Solution of $\min \|Ax - b\|_2$, subject to $x \in \text{span}(V_1 \dots V_K)$ is

$$x^{(K)} = \sum_{k=1}^K t_k V_k.$$

Accuracy issues with the SVD

- ▶ Sometimes: accuracy problems with Matlab.
- ▶ Criterion: compute $[V \ S \ \tilde{V}] = \text{svd}(M)$; \rightsquigarrow check $V \simeq \tilde{V}$.

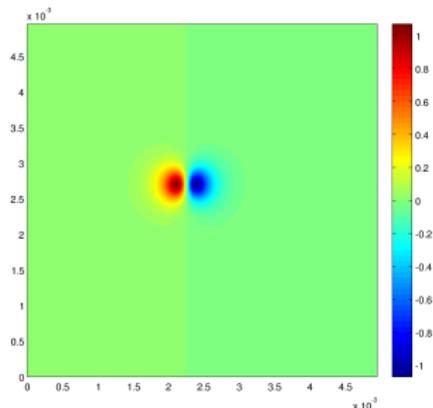


Number of matching bits for each elements of V and \tilde{V}

- ▶ Other criterion: check that $\|b - \sum_{k=1}^K t_k V_k\|$ is decreasing with K .

Using sparse representation?

- ▶ Alternative way to save memory consumption: use sparsity.
- ▶ Column of A : field generated by a single dipole.
 ~→ should decrease quickly.

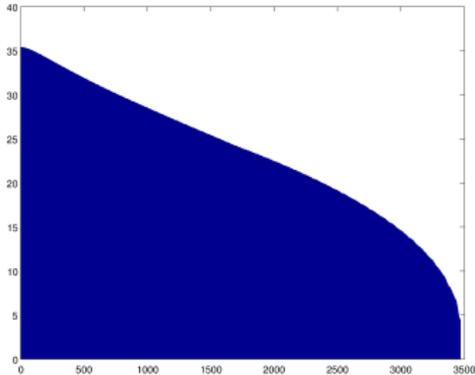


Field produced by a single dipole

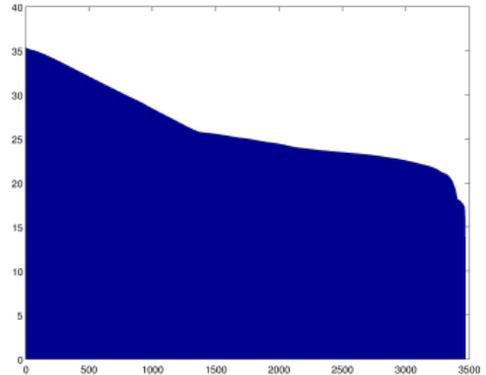
- ▶ Approximate A by replacing values smaller than some threshold by 0.

Sparse representation

- ▶ Sparse representation of $A \rightsquigarrow$ easily fits in memory.
- ▶ Fast computation of A^*A .
- ▶ But does not reduce computation time for the SVD.
- ▶ And... deeply change the behavior of A even with small threshold.



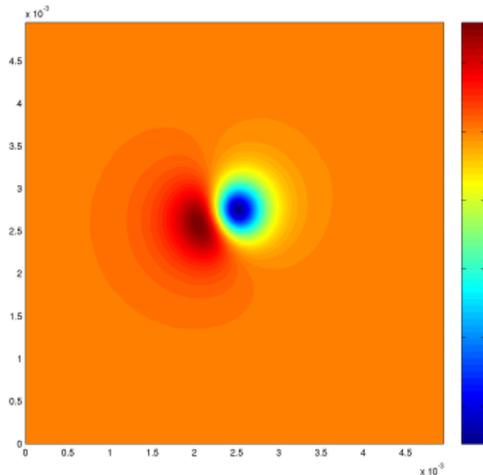
Singular values of A^*A



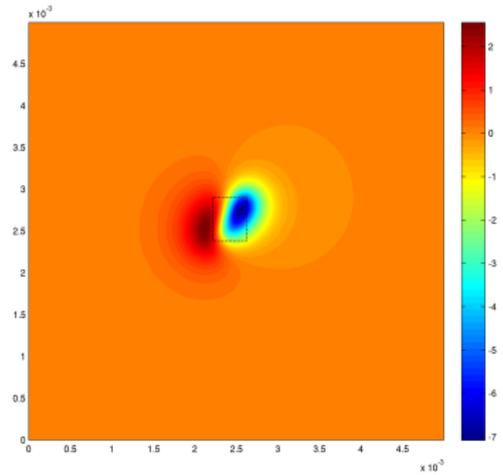
Idem sparse case: threshold = 0.1%

Lunar spherule and synthetic example

- ▶ Lunar spherule: very small sphere of rock.
- ▶ Synthetic example: looks like the spherule example.
 ↪ Allows for benchmarking.



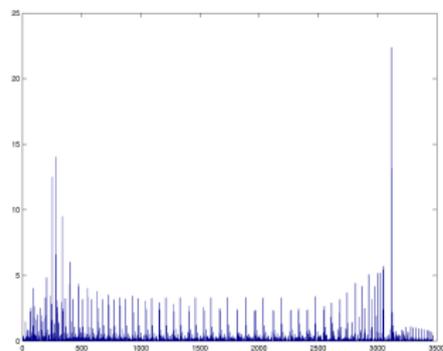
Field measured for the lunar spherule.



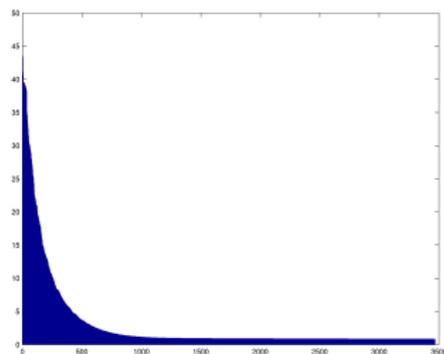
Field of the synthetic example.

Synthetic example, 1st step

- ▶ Magnetization grid: 34×34 , same dimensions as measurement grid.
- ▶ Field very well reconstructed: $\|Ax - b\|_2 / \|b\|_2 \simeq 1.7\%$.
- ▶ But last singular vectors have significant moments.



Amplitude of the net moment
 of V_K



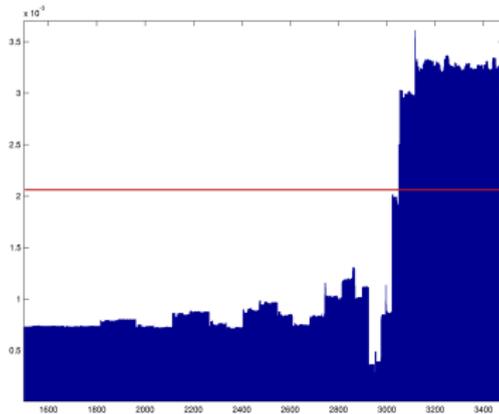
$\|Ax^{(K)} - b\|$.

$$x^{(K)} = \sum_{k=1}^K t_k V_k.$$

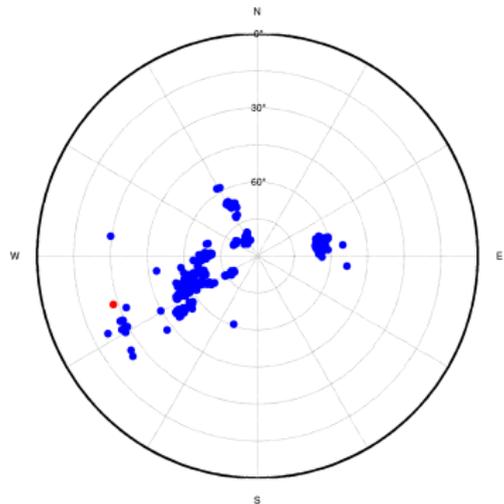
\rightsquigarrow does not permit to deduce the net moment.

1st step: recovered moment

Magnetizations explaining equally well the measured field.
 (true moment in red)



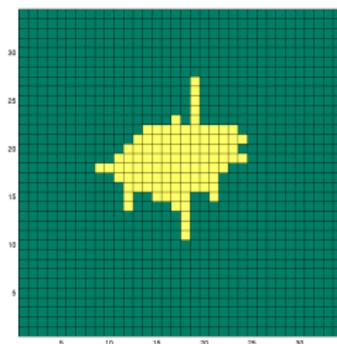
Amplitude of the net moment of $x^{(K)}$.



Direction of the net moment of $x^{(K)}$.

Synthetic example, 2nd step

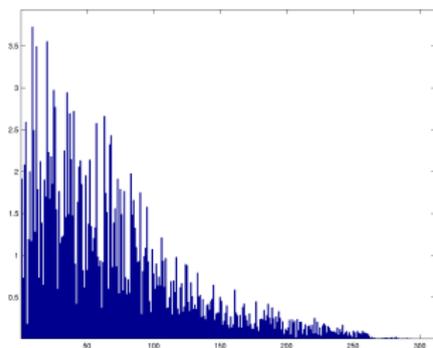
- ▶ We expect the support of the magnetisation to be very localised.
- ▶ We discard points of the magnetization grid, using a thresholding strategy based on the result of the 1st step.
- ▶ Second step with new support.



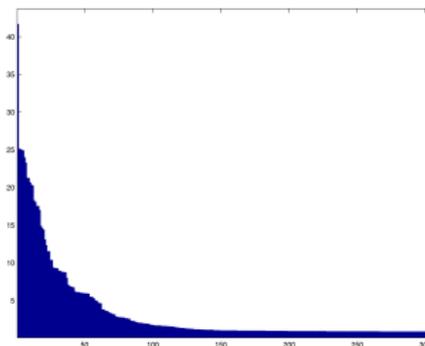
New support in yellow (threshold = 10%)

Synthetic example, 2nd step

- ▶ Field still well reconstructed: $\|Ax - b\|_2 / \|b\|_2 \simeq 1.76\%$.
- ▶ This time, moments of the last singular vectors are very small.
- ▶ Last singular vectors have small moments.



Amplitude of the net moment
 of V_K

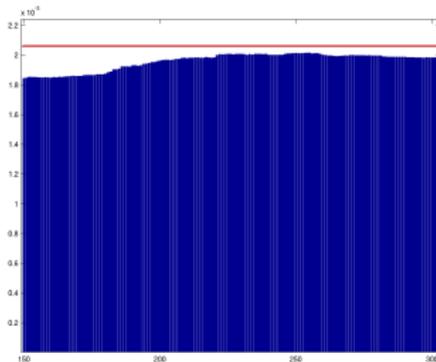


$\|Ax^{(K)} - b\|.$

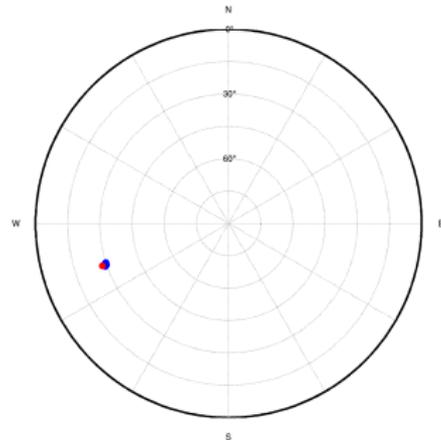
$$x^{(K)} = \sum_{k=1}^K t_k V_k.$$

2nd step: recovered moment

Net moments of the least-square solution on $\text{span}(V_1 \dots V_K)$
 (true moment in red)



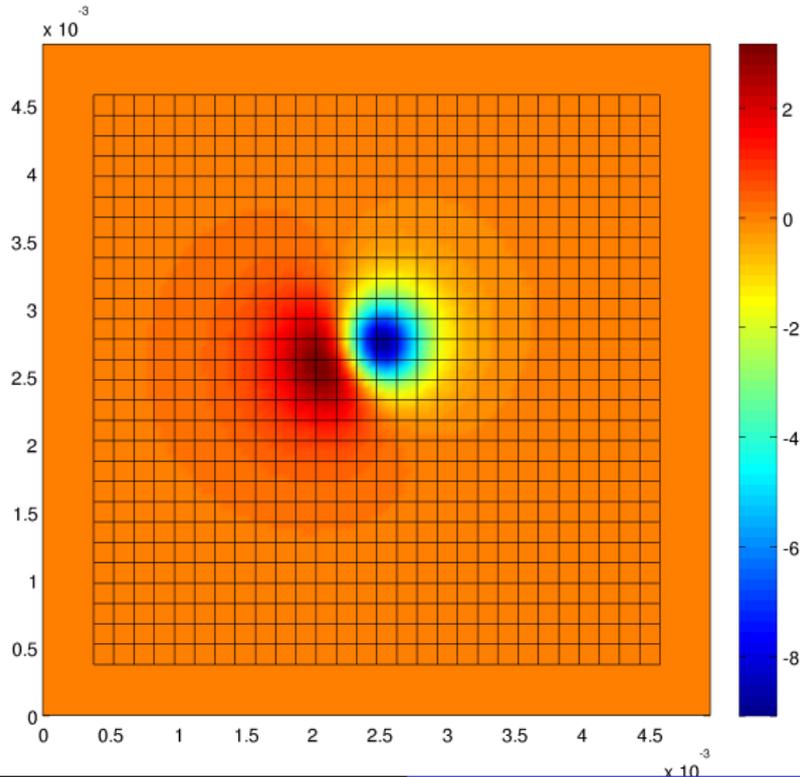
Amplitude of the net moment of $x^{(K)}$.



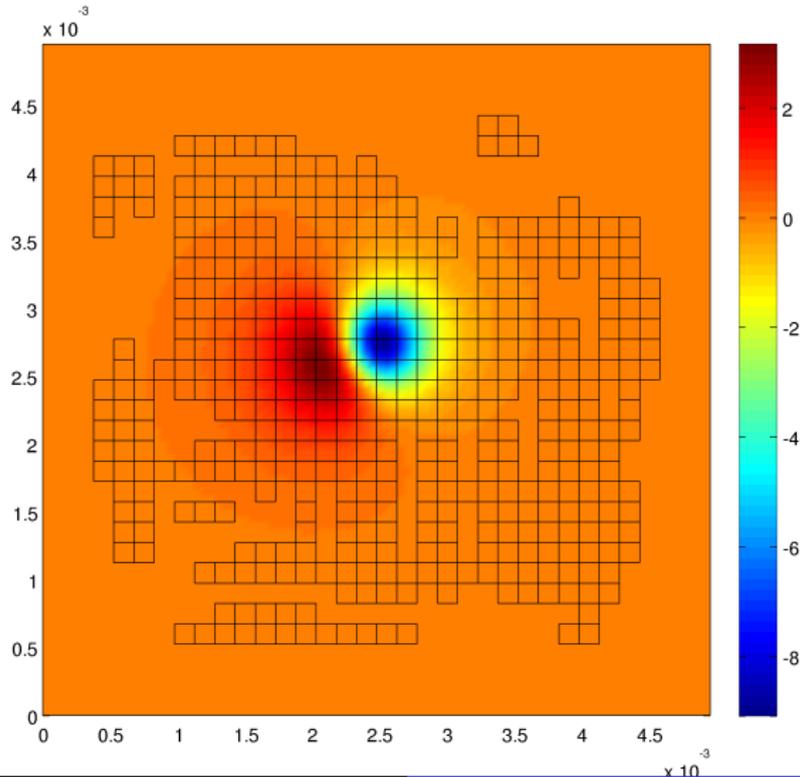
Direction of the net moment of $x^{(K)}$.

- ▶ Net moment remarkably well recovered.
 ~> shrinking the support regularizes the problem.

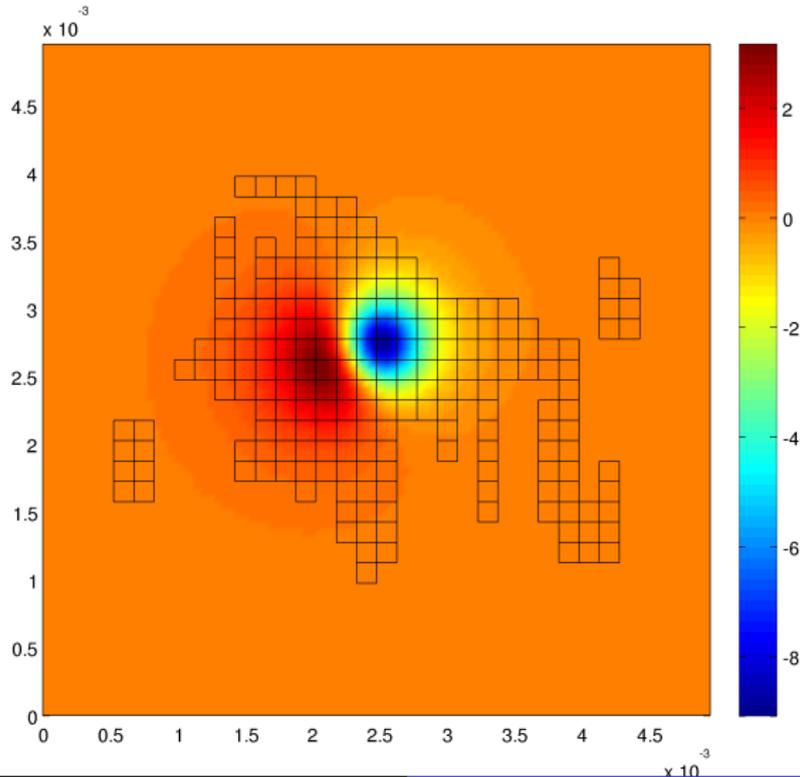
Lunar spherule example



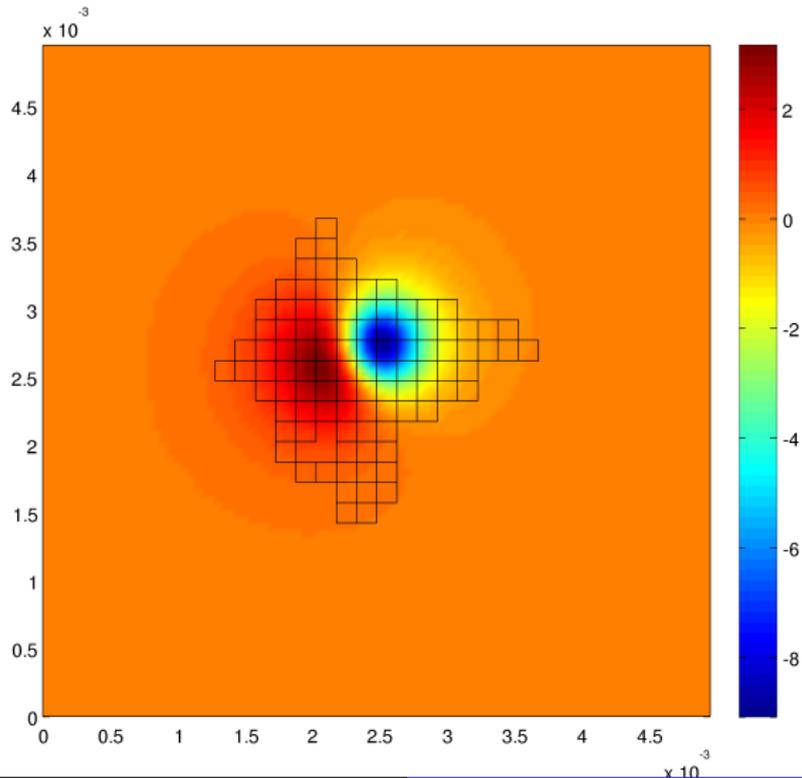
Lunar spherule example



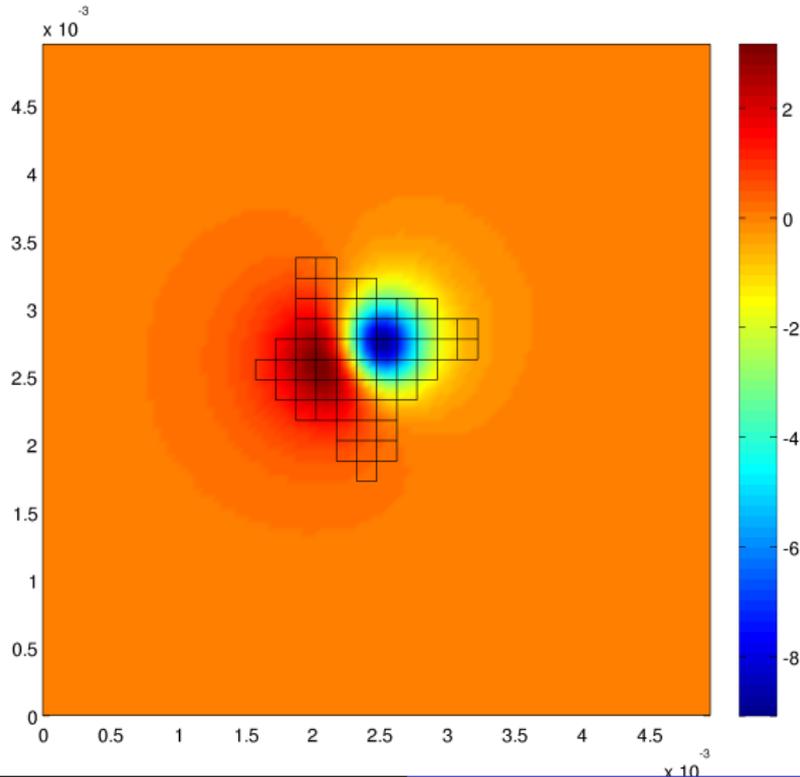
Lonar spherule example



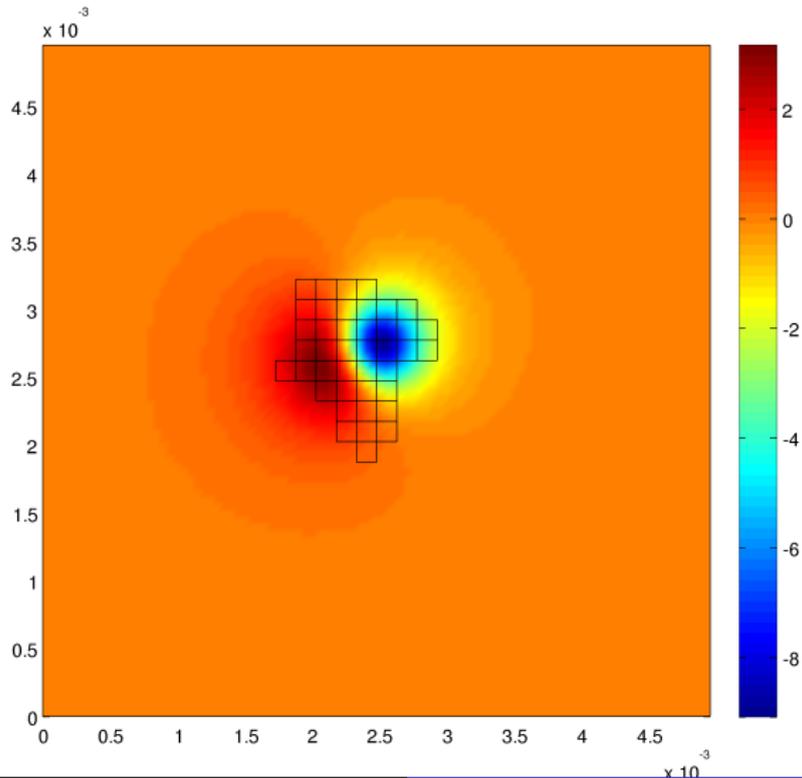
Lonar spherule example



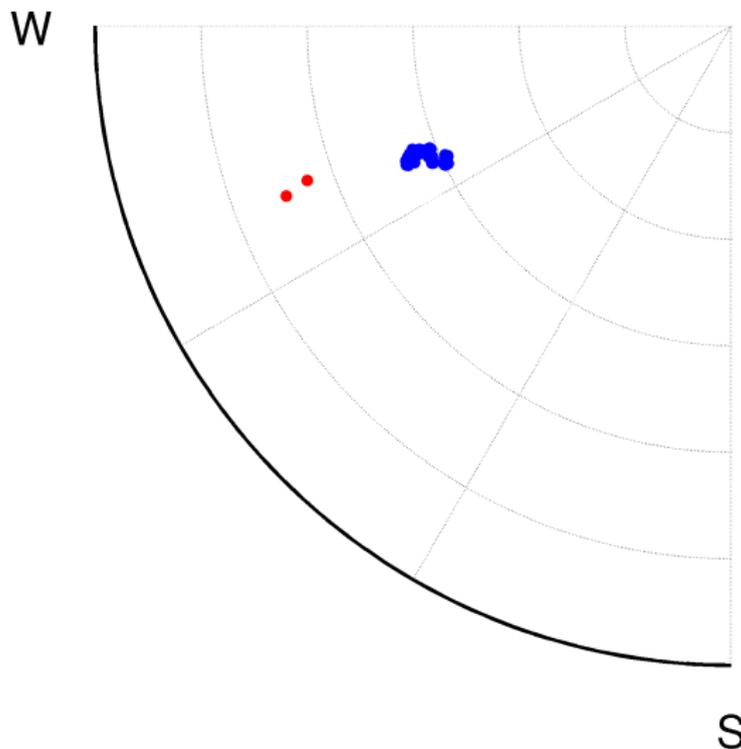
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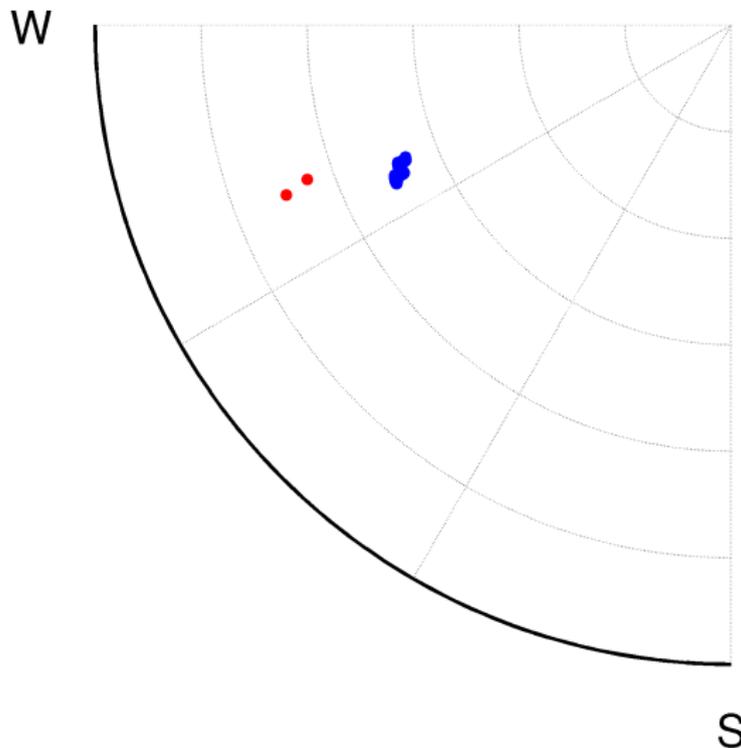
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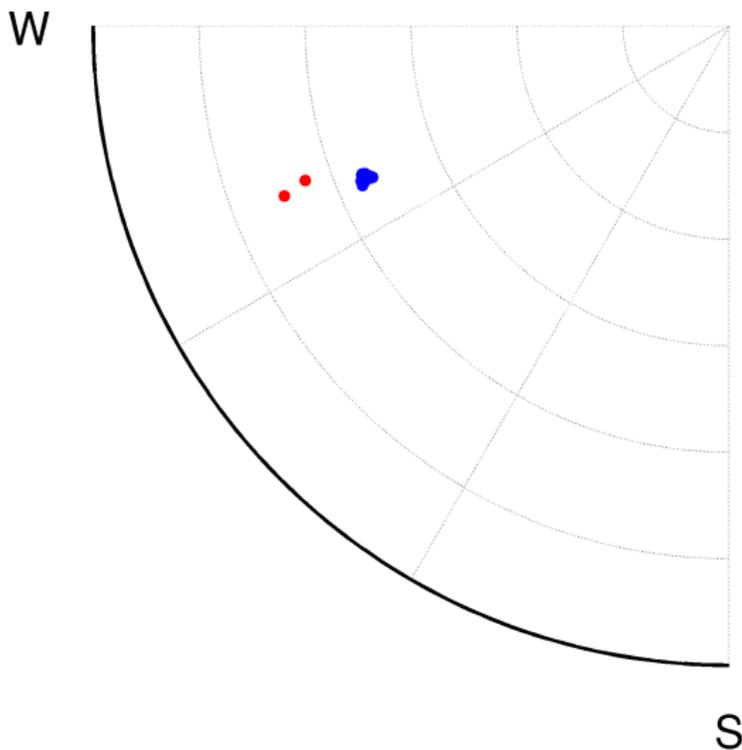
Evolution of the angle of the net moment



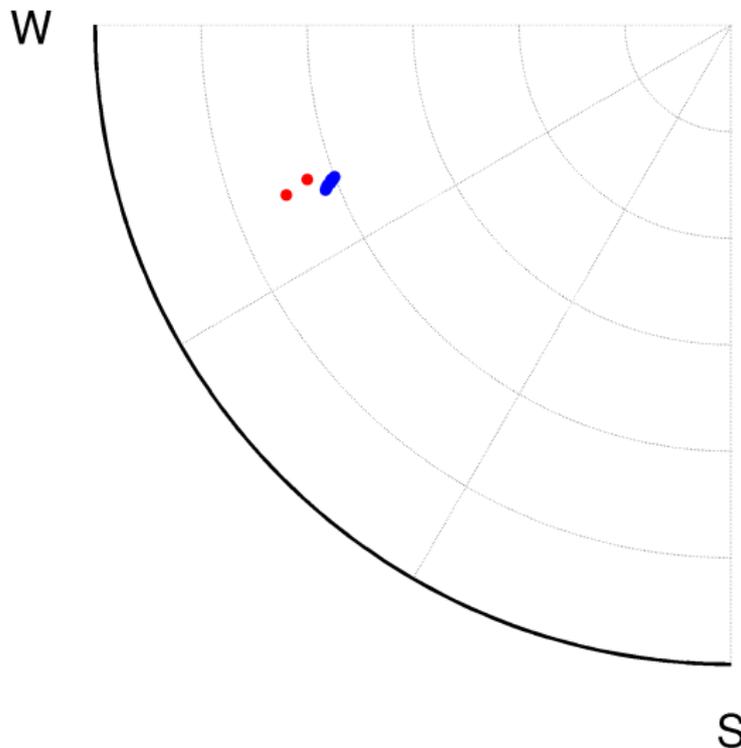
Evolution of the angle of the net moment



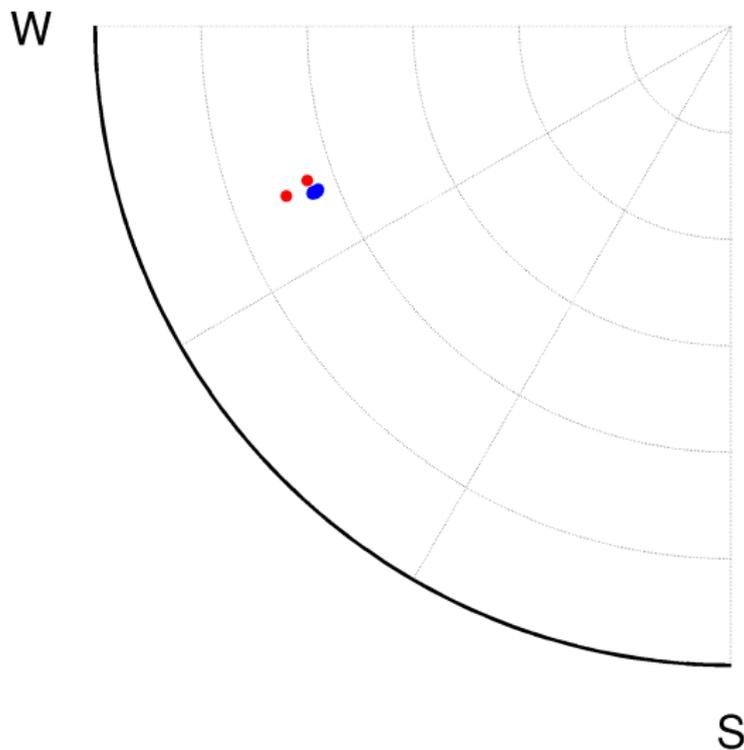
Evolution of the angle of the net moment



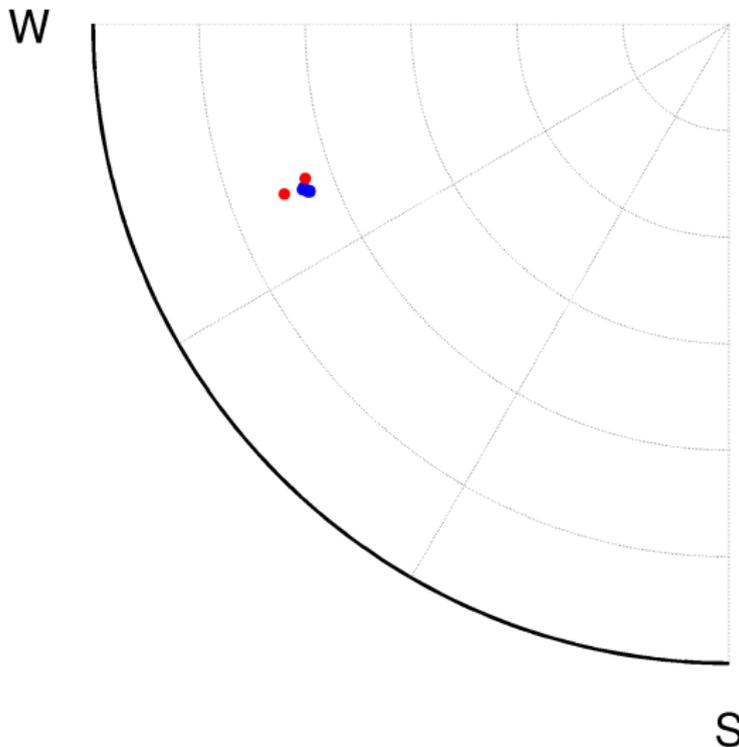
Evolution of the angle of the net moment



Evolution of the angle of the net moment



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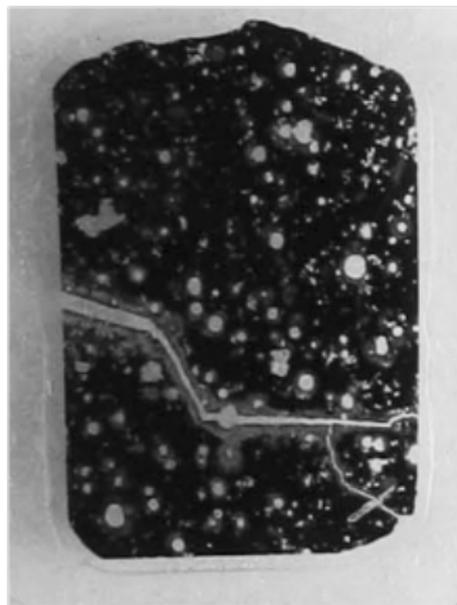


Preliminary results

- ▶ This strategy in two steps has been used on the lonar spherule example and on some chondrules from Allende meteorite.
- ▶ The net moment recovered by this method matches the net moment measured with a magnetometer.

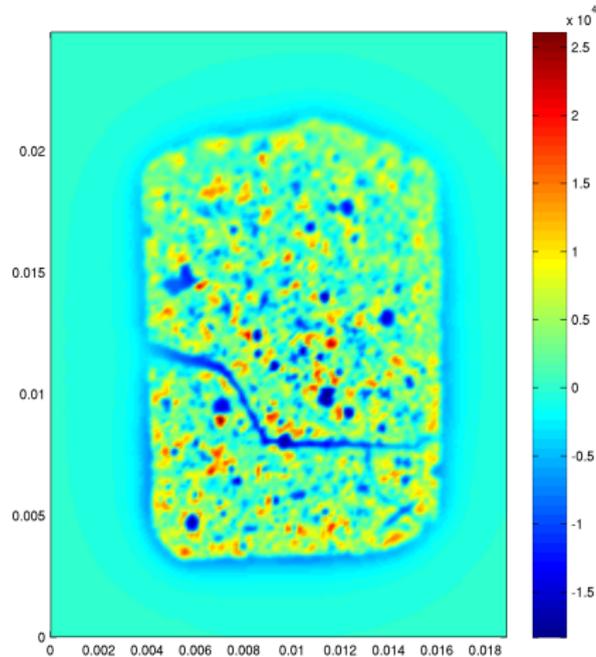
Example	Recovered moment			Measured with magnetometer		
	r	θ	φ	r	θ	φ
Lon. sph.	$5e-6$	117.5	-159.2	$5.31e-6$	112.7	-159.2
A1b1	$1.45e-9$	85.7	11.7	$1.54e-9$	87.2	14.9
A1b4	$7.4e-10$	160	155	$6.28e-10$	155	160
A1b6	$1.8e-11$	92	223.8	$1.73e-11$	93.2	234.5

Non-localised example



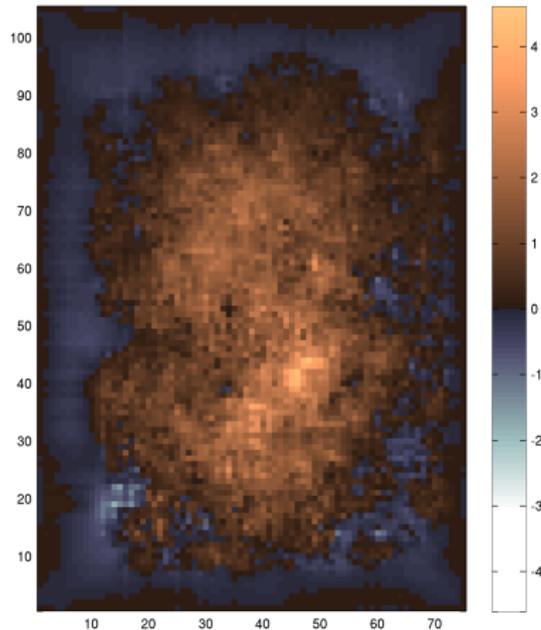
Photograph of an Hawaiian basalt

Non-localised example



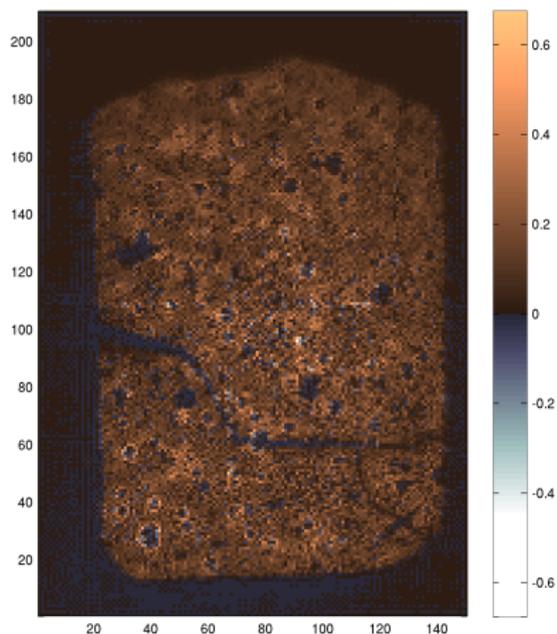
Measured field B_z

Non-localised example



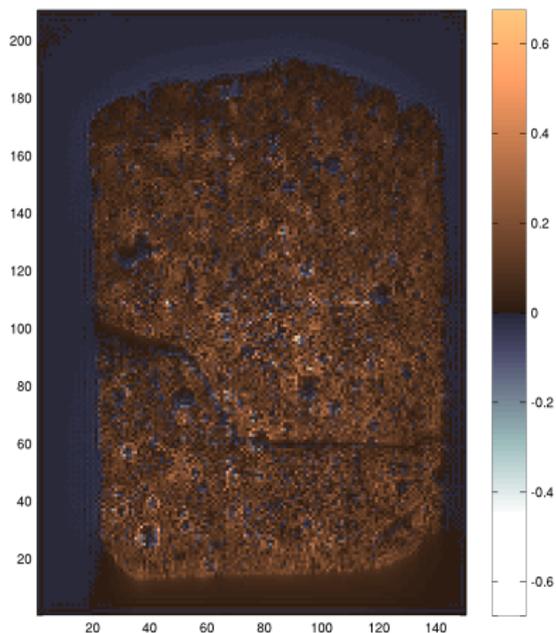
Recovered 3D magnetization (normal component)

Non-localised example



Recovered unidirectional magnetization (right direction)

Non-localised example

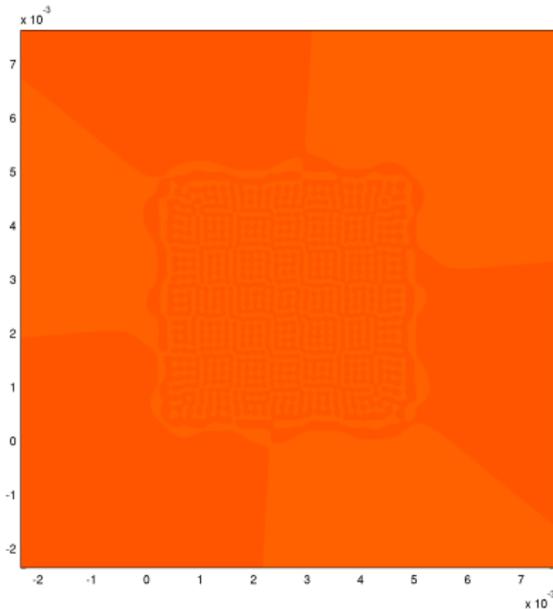


Recovered unidirectional magnetization (wrong direction)

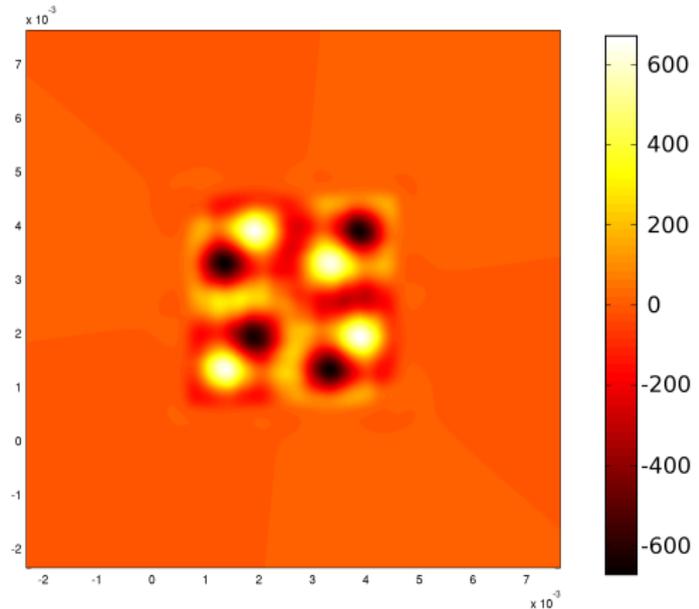
What next?

- ▶ Use less naive strategies to compute the net moment of the magnetization.
- ▶ Once a plausible magnetization is recovered, find an equivalent unidirectional magnetization.
- ▶ Problem: existence of fairly silent magnetization from above, but not from below.
- ▶ Solution? Measure from both sides?

Silent source from only one side



B_z measured at height $h = 270\mu\text{m}$



B_z measured at height $h = -270\mu\text{m}$

Field produced by a 28×28 grid of uniformly magnetized squares.